

Diurnal cycles in thermokarst lakes of a permafrost peatland

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ABSTRACT. Despite the importance of surface waters of permafrost landscapes in carbon (C) emission, dissolved C, metal storage and export, the diel pattern of dissolved components and gases in thermokarst lakes remain virtually unknown. Here we discovered a factor of 2 to 3 higher CO₂ concentrations and fluxes during the night compared to day-time in the high-DOC lake. The emission fluxes in the low-DOC lake increased from zero to negative values during the day to highly positive values during the end of the night and early morning. The bulk of dissolved (< 0.45 μm) hydrochemical parameters remained highly stable with ± 10% variation in concentration over 2 days of observation. Overall, the impact of diel cycle on dissolved CH₄, DOC, nutrient and metal concentration was below 10%. However, neglecting night-time period may underestimate net CO₂ emission by ca. 30 to 50% in small organic-rich thaw ponds and switch the CO₂ exchange from uptake/zero to net emission in larger thermokarst lakes. Given the dominance of large lakes in permafrost regions, the global underestimation of the emission flux may be quite high. As such, monitoring CO₂ concentrations and fluxes in thermokarst lakes during months of extended night time (August to October) is mandatory for assessing the net emissions from lentic waters of frozen peatlands.

Keywords: GHG emission, thermokarst, pond, trace metal, organic carbon, warming

1. Introduction

The emission of greenhouse gases (GHG), particularly CO₂ and CH₄, from inland waters at high latitudes is a critical issue of these aquatic system behaviors under on-going climate warming (Serikova et al., 2019). The majority of dystrophic to mesotrophic aquatic systems in frozen peatlands drastically differ from oligotrophic northern rivers and lakes in the sense that the peatland thermokarst lakes and ponds are i) highly humic, and typically contain between 10 and 100 mg L⁻¹ of DOC and ii) quite acidic, with pH ranging between 3.5 and 5.5 (Shirokova et al., 2019). The working hypotheses of this study is that the variability in external factors like temperature and light will exert primary control on DOC and related trace elements, and this will produce a diel cycle in C and elements concentrations and GHG emissions in thermokarst lakes. In order to test these hypotheses, we studied

two thaw ponds with different size and DOM content and we quantified (1) the variation of dissolved C and metal concentration between day and night and (2) the degree of diel variation in CO₂ flux, in order to reveal the control of hydrochemical parameters by external factors such as temperature and light.

2. Material and methods

We studied typical thermokarst lake (Lake Trisino, S_{area} = 16920 m²; S_{watershed} = 0.108 km², average depth 0.5 m) and thaw pond (Chernoe, S_{area} = 1830 m²; S_{watershed} ≤ 500 m², average depth 0.3 m) located in the discontinuous permafrost zone of the Western Siberian Lowland (WSL). Observations were performed during 48 h on July 26-28th 2018, at stable anticyclone conditions without any precipitation event. Every 2 hours we measured the dissolved oxygen, specific conductivity, pH, water temperature, partial pressure

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of CO₂ (pCO₂), dissolved CH₄, CO₂ fluxes (FCO₂), DOC, DIC, and dissolved major and trace elements, chloride, sulfide, fluoride and carboxylic organic ligands (lactate, acetate, propionate, formate, butyrate, pyruvate, galacturonate, glutarate, malate, tartrate, oxalate, phthalate, quinate, and citrate) and total microbial cell concentration.

3. Results and Discussion

The water temperature of the two lakes followed a similar diurnal cycle with a minimum temperature reached approximately 2 h after the sunrise. The concentration of DOC, phosphate, silica and carboxylate remained constant (formate, oxalate, pyruvate) or fluctuated without any diel pattern (acetate and lactate). The total cell number and biomass of bacteria, DIC, Cl⁻, SO₄²⁻, most major and trace elements did not exhibit any sizeable variations over the diurnal cycle as their concentrations remained stable (within ± 10-15%). The pCO₂ demonstrated clear diurnal pattern with a factor of 2 to 4 times higher concentrations achieved at the end of the night and early morning compared with the middle of the day. The FCO₂ was almost nil or negative during the day time in Trisino (-0.02 ± 0.02 g C m⁻² d⁻¹), but positive in Chernoe (0.12 ± 0.07 g C m⁻² d⁻¹). During the night, the FCO₂ strongly increased and became positive in Trisino (0.12 ± 0.1 g C m⁻² d⁻¹), whereas in Chernoe, the night maximum was 2 to 5 times higher than the day minimum. Overall the impact of diel cycle on dissolved CH₄, DOC, nutrient and metal concentration was below 10%. However, neglecting night-time period may underestimate net CO₂ emission by 30 to 50% in small organic-rich thaw ponds and switch the CO₂ exchange from uptake/nil to a net emission in larger thermokarst lakes.

4. Conclusion

We observed stable pH, SUVA₂₅₄ and concentrations of DOC, dissolved CH₄, nutrients,

carboxylic acids, most major and trace elements in waters of acidic and humic thermokarst lake and thaw pond, during a continuous 2-days observation at generally anticyclone weather. In contrast, the CO₂ concentration and emission exhibited clear diurnal pattern with a maximum after the sunrise (minimal surface water temperature) and a minimum during late afternoon day-time (maximum surface water temperature). It is possible that the balance between benthic respiration of dissolved and particulate organic matter and primary productivity of periphyton, aquatic macrophytes and submerged mosses was responsible for enhanced CO₂ emission during the night and CO₂ decrease or even uptake during the day. Neglecting night-time CO₂ flux, especially in large thermokarst lakes, can sizably underestimate the region-upscaled global C emission. This result convinces the need to study diurnal variations in C emission of large (10-500 ha) thermokarst lakes as they mostly contribute to overall inland water surface of the region.

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